

Late Quaternary Paleoenvironmental Dynamics on the Cameroonian Continental Shelf (Gulf of Guinea): Palynological and Sedimentary Insights

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Abstract

The palynological study of core C61, retrieved from the Cameroonian shelf near the Sanaga Delta, offers new insights into the evolution of coastal environments and vegetation dynamics in Central Africa since the end of the Last Glacial Maximum. This 6-meter-long marine sequence spans the transition from the Late Pleistocene to the Late Holocene. Pollen analysis of 26 stratigraphic levels revealed rich and diverse assemblages (224 taxa, 93 families), allowing the identification of three major paleoecological phases: 1) a warm and humid period dominated by shrub savannas (*Cyperaceae*, *Bridelia*, *Combretaceae*) inland and *Rhizophora* mangroves along the coast; 2) a more open landscape phase marked by the expansion of grasses (*Poaceae*), *Commelina*, and spores, though without strong evidence of aridification; 3) a phase characterized by alternating humid and drier intervals, reflected in the decline of herbaceous taxa and the development of forest species (*Caesalpiniaceae*, *Sapotaceae*, *Sacoglottis*, *Podocarpus*). These stages roughly correspond to the Early, Middle, and Late Holocene. While the sequence captures broad environmental trends, gaps in chronological resolution limit the detection of short-term climatic events such as the ca. 2.5 ka dry phase observed in nearby lake archives. The C61 record highlights the high sensitivity of coastal vegetation in Central Africa to Holocene climatic variability, particularly to latitudinal shifts of the Intertropical Convergence Zone (ITCZ). By linking marine pollen signals to regional vegetation dynamics, this study contributes to refining our understanding of the long-term ecological responses of tropical coastal systems to climate forcing.

Keywords

Paleoenvironments, Climate, C61 Marine Core, Pollen, Central Africa, Cameroon, Holocene

1. Introduction

Each year, the Conferences of the Parties (COP) underscore the profound impact of human activities on the environment, particularly through fossil fuel exploitation, urbanization, and deforestation (IPCC, 2021; Malhi et al., 2014). These processes severely disrupt ecosystems and contribute to global climate instability. However, in the context of this study—which focuses on Holocene and Late Pleistocene periods—such anthropogenic impacts are not relevant, as the recorded environmental changes occurred well before significant human intervention.

In this context, the study of natural archives is crucial for improving our understanding of past interactions between climate and vegetation. Paleoenvironmental analyses, especially those based on pollen records, make it possible to compare past and present climatic conditions. In the case of core C61, the focus is not on anticipating future climate change but rather on documenting past environmental dynamics to establish robust paleoecological baselines (Lézine et al., 2009; Dupont, 2011). The palynological approach has been widely employed in Africa to reconstruct past environmental and climatic conditions, both in terrestrial and marine contexts.

In West Africa, numerous studies have examined both continental sedimentary sequences (Maley, 1987; Lézine, 1988; Lézine & Cazet, 2005; Assi-Kaudjhis et al., 2010) and marine sequences (Hooghiemstra et al., 1986; Caratini et al., 1987; Fredoux, 1980; Fredoux, 1994; Marret, 1994; Dupont & Weinelt, 1996; Dupont, 2012). However, detailed knowledge of hydrodynamic processes, the origin of pollen grains, and their modes of transport and deposition within sedimentary basins remains limited (Vincens et al., 1994; Lézine & Casanova, 1989). Recent progress has been achieved through studies of modern pollen dispersal on the Cameroonian continental shelf, which have highlighted the crucial role of marine hydrodynamics in shaping the spatial distribution of pollen grains (Bengo et al., 2025).

In Central Africa, most research has focused on lake sediment sequences (Brenac, 1988; Elenga et al., 1992, 1994, 2001, 2004; Reynaud-Farrera, 1995; Reynaud-Farrera et al., 1996; Stager & Anfang-Sutter, 1999), which provide valuable pollen records for reconstructing local paleoecosystems. These studies have significantly enhanced our understanding of the evolution of equatorial forests and their sensitivity to climatic variability (Ngomanda et al., 2005, 2009a, 2009b; Lebamba et al., 2009, 2012; Maley et al., 2017). However, marine pollen records remain scarce in this region, particularly off the Gulf of Guinea and along the Cameroonian continental shelf (Bengo et al., 2025; Van Campo & Bengo, 2004).

Yet, these marine deposits represent a crucial source of information, as they integrate a regional signal across the Gulf of Guinea (Lopez-Merino et al., 2018; Julier et al., 2018; Hernandez et al., 2021; Tahi, 2022), Central Africa (Bengo & Maley, 1991; Roche, 1991; Runge, 2007), and Eastern and Southern Africa (Dupont et al., 2022; Neumann et al., 2025; Yao et al., 2025).

This local and regional information can be compared with marine and deltaic records from similar tropical environments—for example, offshore the Congo River, where deep-sea fan deposits integrate continental inputs and reflect regional hydro-sedimentary dynamics (Dupont & Agwu, 1991; Giresse et al., 1995); in the Niger Delta, where palynological and sedimentological data document the evolution of fluvio-marine environments in response to climatic and eustatic variations (Giresse et al., 1995); along the Ogooué margin, which provides an integrated 26,000-year record of terrestrial and marine environmental changes off Gabon (Ngomanda et al., 2005; Gasse, 2000); and even along the Amazon margin, where oceanic sediments reveal fluctuations in hydrological and sedimentary inputs linked to tropical climate variability (Maley & Brenac, 1998). Such comparisons provide a better understanding of the sedimentary and ecological dynamics of tropical fluvio-marine systems.

It is within this framework that the present study was undertaken, based on the analysis of core C61, collected from the Cameroonian continental shelf during the oceanographic campaigns CAMPUS-Cameroon, CHRIS-Elf, and ECOFIT (CNRS & ORSTOM). Several datasets from these missions have already been analyzed, particularly focusing on the mineral (Ngueutchoua, 1996) and pollen (Bengo, 1996) fractions of the sediments. These studies identified distinct lithological units and sedimentary structures that were interpreted in relation to the regional Holocene transgression and associated environmental changes (Ngueutchoua, 1996; Giresse et al., 1995), thus providing a stratigraphic and paleoenvironmental framework that underpins the present investigation. However, their potential for integration into a regional paleoclimatic reconstruction remains underexplored.

The main objectives of this study are

- to compare present-day flora with that of past periods in order to highlight major trends in vegetation evolution;
- to reconstruct past environments and climates based on pollen analysis;
- and to identify the main ecological transitions through time by analyzing the variability of key environmental indicators, particularly during the Holocene.

A more specific assessment of possible anthropogenic influences is not attempted here, as this would require a focused analysis of recent intervals (post-2500 or post-1000 BP), which falls outside the scope of the present study.

This study, therefore, aims to enhance our understanding of climate-vegetation interactions in Central Africa by providing a robust reference framework for interpreting ecosystem responses to past and present environmental changes. While future dynamics remain uncertain and cannot be directly inferred from the available data, the paleoclimatic record from core C61 offers valuable insights into the

range of natural variability and the resilience of ecosystems to climate fluctuations.

2. Materials and Methods

2.1. Environmental Framework

2.1.1. Location and General Context of the Study Area

Core C61 was collected in the Gulf of Guinea, on the central part of the Cameroonian continental shelf, at a depth of approximately 200 meters (**Figure 1**). This area lies along the upper margin of the continental slope and is strongly influenced by riverine inputs from the Sanaga River, as well as other coastal rivers such as the Wouri, Nyong, and Ntem. These rivers transport fine-grained sediments rich in organic matter and pollen into the marine environment. The Sanaga River watershed, which spans roughly 133,000 km², covers a broad ecological gradient from humid equatorial forests in the south to drier savanna regions in the center and north of Cameroon (**Bengo et al., 2020**). This wide range of source environments contributes to the mixed continental signature of the sedimentary inputs observed offshore. Since most detrital particles reaching the shelf originate from continental erosion, the marine sedimentation in this area primarily reflects terrigenous input, aside from biogenic components produced in situ.

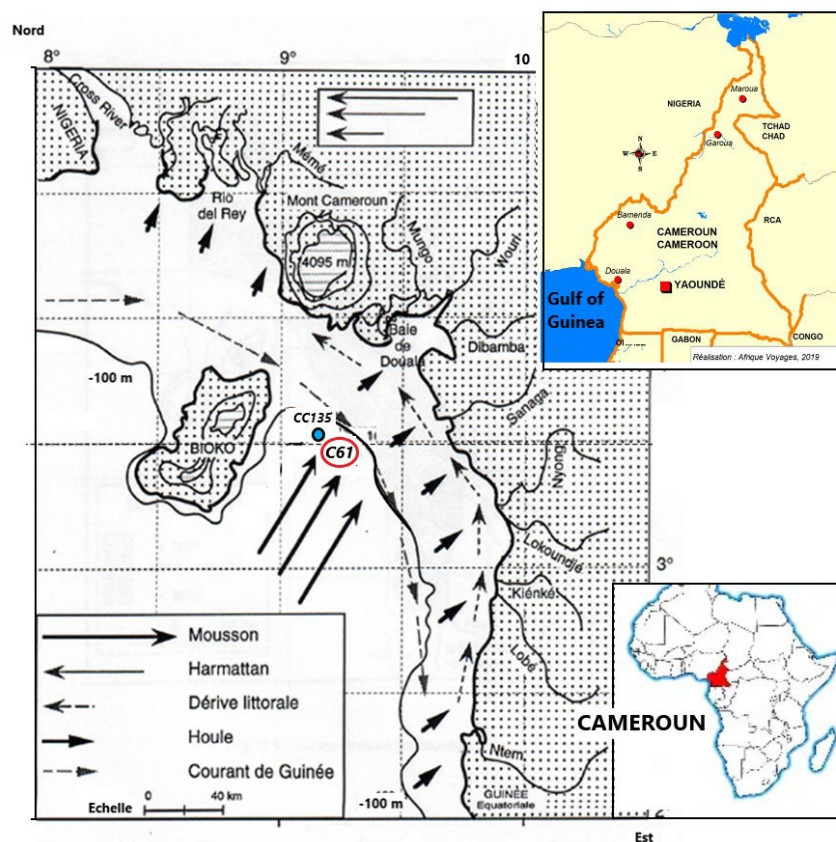


Figure 1. Location of the sampling sites on the Cameroonian continental shelf.

The regional oceanic circulation, dominated by the Equatorial Countercurrent and the Guinea Current (Bourlès et al., 1999; Braga et al., 2004), facilitates the dispersion of particles and the development of a continuous sedimentary record that accurately reflects paleoenvironmental variations from the Late Pleistocene to the recent Holocene (Giresse et al., 1996; Bengo et al., 2025).

2.1.2. Geological and Geomorphological Framework

The Cameroonian continental shelf is a geological structure inherited from the opening of the South Atlantic during the Early Cretaceous, characterized by N60-oriented faults that produced an alternation of horst and graben-type tectonic blocks (Aloisi et al., 1995). The sedimentary cover mainly consists of clayey and silty deposits of fluvial origin, largely supplied by the Sanaga River basin. The continental relief, organized in successive steps from the coastal plain to the Adamawa Plateau, promotes intense erosion and a rapid transfer of terrigenous material to the sea. In marine settings, sedimentary dynamics are expressed by recent deltaic deposits in the north, whereas in the south, relict sediments such as glauconitic and bioclastic sands—remnants of the Last Glacial Maximum and subsequent Holocene transgressions—predominate (Giresse et al., 1995; Nguetchoua, 1996). Overall, this region represents a complex sedimentary system shaped by the combined influence of tectonic, climatic, and hydrological processes, which together account for the high heterogeneity of the deposits.

2.1.3. Pedological Framework of the Sanaga River Basin

The Sanaga River basin, the main source of sediments supplying the Cameroonian continental shelf, exhibits a wide diversity of soils that directly influence the composition of particles exported to the marine environment. In the humid equatorial zone of southern Cameroon, kaolinite-rich ferrallitic soils, developed on ancient Precambrian basement rocks, are predominant. Swampy areas host hydromorphic soils enriched in organic matter, while the volcanic regions of the west and southwest are covered by Andosols derived from basic volcanic rocks (Segalen, 1967; Martin, 1966). Under the combined effects of heavy rainfall and stepped topography, these soil formations undergo intense erosion, resulting in the transport of clays, silts, and pollen grains to the ocean.

2.1.4. Atmospheric and Oceanographic settings

The climate of the Cameroonian coast is largely controlled by the Intertropical Convergence Zone (ITCZ), whose seasonal movements determine an alternation between a wet season dominated by oceanic air masses and a dry season marked by the continental Harmattan winds (Leroux, 1983; Piton, 1987; Suchel, 1988). These atmospheric variations directly influence regional ocean circulation in the Gulf of Guinea (Bourlès et al., 1999; Braga et al., 2004; Kolodziejczyk, 2008), characterized by a northward coastal drift current and the southward-flowing Guinea Current (Figure 1). At the surface, these currents facilitate the transport and dispersion of fine particles and pollen material along the coast, while in deeper wa-

ters, the Guinea Current promotes sediment redistribution offshore.

2.1.5. Current Climate and Vegetation Formations

The coastal region of Cameroon experiences a humid equatorial climate, characterized by mean annual temperatures ranging from 27 to 29°C and exceptionally high rainfall, locally exceeding 10,000 mm per year at the foot of Mount Cameroon (Suchel, 1988). This high level of humidity promotes the development of dense and diverse vegetation, dominated by humid evergreen forests rich in Caesalpiniaceae, Euphorbiaceae, and Ulmaceae (Letouzey, 1968; White, 1983). Along the margins of flooded and lagoonal areas, swamp forests and mangroves dominated by *Rhizophora* and *Avicennia* thrive, whereas inland from the coast, semi-deciduous forests give way to wooded savannas in drier regions. This vegetation mosaic, governed by climatic and topographic gradients, illustrates the gradual transition between forest and savanna biomes. Its present-day floristic composition is reflected in the pollen signal preserved in marine sediments, thereby providing an essential reference framework for the reconstruction of Holocene paleoenvironments (Maley, 1987; Bengo et al., 2020; Bengo et al., 2025).

2.2. Methods

2.2.1. Sampling Protocol

A total of 26 samples were collected: 16 at 10 cm intervals in the upper section (0 - 150 cm), where sedimentation was relatively slow, and 10 at 30 cm intervals in the lower section, characterized by more rapid sedimentation. This sampling strategy provides representative coverage of paleoenvironmental variations spanning from the end of the Pleistocene to the Late Holocene.

2.2.2. Palynological Processing

The palynological processing of sediments from core C61 was conducted following the standard procedures described by Faegri and Iversen (1992). Each sample underwent a series of chemical treatments to remove carbonates, silicates, and non-pollen organic matter, followed by repeated washing and centrifugation to obtain clean residues. The resulting fractions were then mounted on microscope slides using glycerine jelly as the mounting medium. This procedure ensured good preservation and optimal visibility of pollen grains and spores for both quantitative and qualitative analyses.

2.2.3. Pollen Observation, Identification, and Classification

Observations were performed using an optical microscope at magnifications ranging from $\times 400$ to $\times 1000$. For each sample, a minimum of 300 pollen and spore grains were counted to ensure adequate statistical representativeness. Taxonomic identification was based on published atlases (Maley, 1970; Bonnefille & Riollet, 1980; Salard-Cheboldaëff, 1980, 1981, 1982, 1983) and on the reference collection of the Palynology Laboratory at the University of Montpellier 2. The identified taxa were subsequently grouped according to their ecological affinities: dense ev-

ergreen forests, semi-deciduous forests, savannas, swamp formations, mangroves, and aquatic plants. This classification made it possible to establish ecological indicator groups used for the reconstruction of paleoenvironmental and paleoclimatic conditions.

2.2.4. Ecological Significance and Chronological Framework

The interpretation of pollen assemblages is based on the identification of bioindicator taxa that reflect climatic and ecological conditions through time. *Poaceae* and *Bridelia* indicate dry phases dominated by savanna formations, whereas *Caesalpinaceae*, *Sapotaceae*, and *Ulmaceae* point to humid periods favorable for the development of dense forest vegetation. The presence of *Podocarpus* is associated with cooler conditions at higher elevations or during episodes of climatic cooling. When considered within the context of the Late Glacial and Holocene, these indicators allow the establishment of a relative chronology of paleoenvironmental changes and enable the correlation of regional vegetation dynamics with major global climatic fluctuations.

3. Results

3.1. Litho-Sedimentary Characteristics and Chronology of Core C61

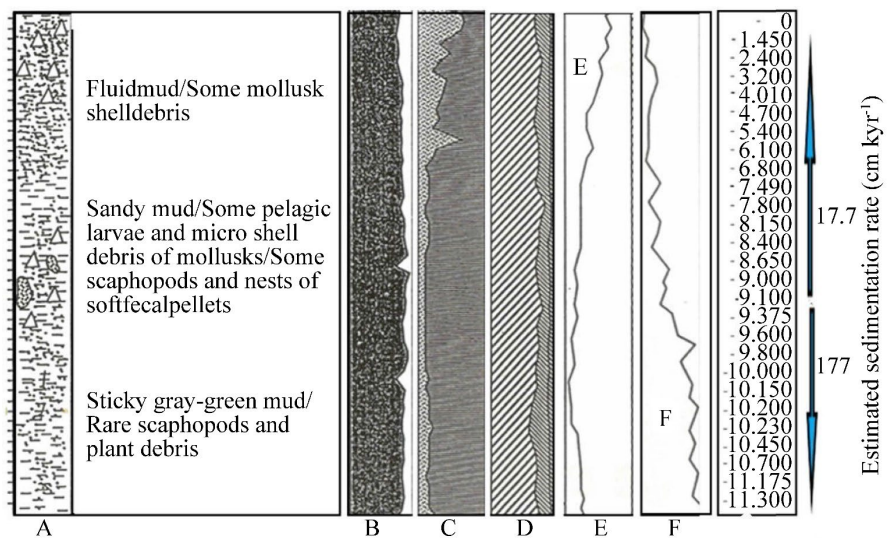
Geochemical analyses of core C61 show a progressive decrease in CaCO_3 content toward the top of the sequence, while total organic carbon increases. This trend is interpreted as a shift toward enhanced marine productivity and environmental stabilization in the uppermost levels. The clay mineral assemblage is dominated by kaolinite, with subordinate amounts of smectite and illite, indicating a predominantly continental origin of detrital inputs, primarily from the Sanaga River basin.

The lithological profile of the core, based on previous studies by [Ngueutchoua \(1996\)](#) and [Giresse and Ngueutchoua \(1998\)](#) ([Figure 2](#)), reveals three main sedimentary units. The basal unit (540 - 355 cm) consists of grey-green clayey mud enriched in plant debris, reflecting strong fluvial input and rapid deposition. The middle section (355 - 145 cm) is composed of sandy mud containing micro-shell fragments, suggesting a transitional depositional environment with increasing marine influence. The uppermost unit (145 - 0 cm) is made up of soft, homogeneous grey mud rich in mollusk remains, characteristic of a low-energy, stabilized marine setting.

To establish the chronological framework of the sequence, six levels of core C61 were dated by radiocarbon (^{14}C) analysis performed on both bulk organic matter and mollusk shells ([Ngueutchoua, 1996](#); [Giresse et al., 1995](#)). The resulting calibrated ages (cal yr BP) range from approximately 11,330 cal yr BP at the base to 1450 cal yr BP at the top of the core.

The sedimentation rate was not uniform throughout the sequence. Between 11,300 and 9100 cal yr BP, the average sedimentation rate was approximately 157

cm per 1000 years. After 9100 cal yr BP, the rate dropped sharply to about 24 cm per 1000 years, and during the last millennium, it decreased further to approximately 10 cm per 1000 years (Figure 3).



A—Mollusk shell debris and Nests of soft fecal pellets; B—% Water; C—% Sand and % Pelites; D—% Kaolinite + % Smectite + % Illite; E—% CaCO₃; F—% Organic carbon. C¹⁴ dating (0 - 11,300). Estimated sedimentation rate (cm/10³ years).

Figure 2. Lithological profile of core C61 (Ngueutchoua, 1996).

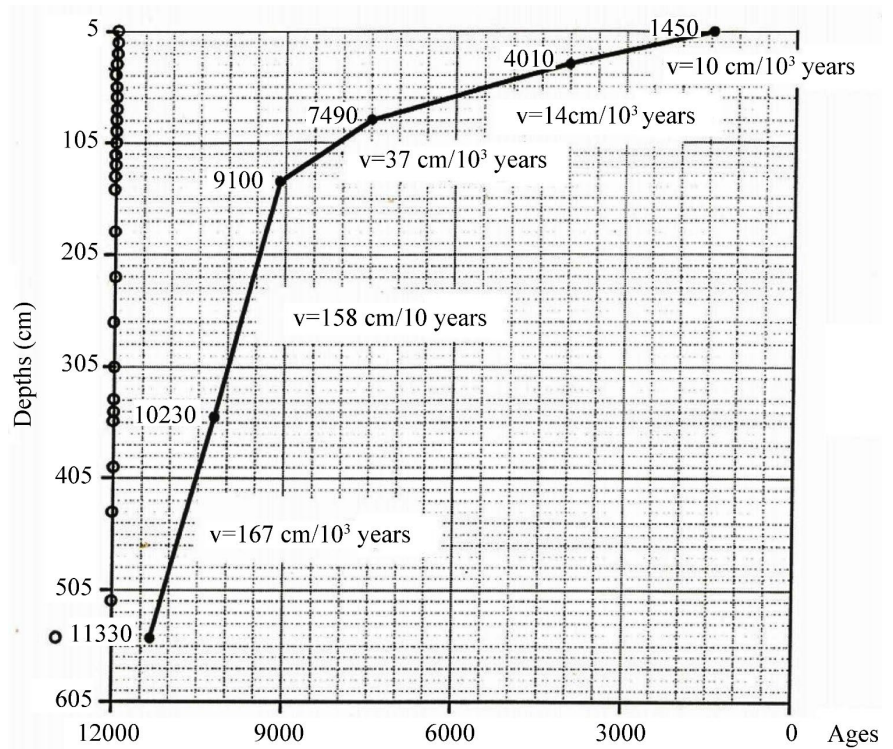


Figure 3. Age-depth model and sedimentation rates derived from calibrated ¹⁴C dates (Ngueutchoua, 1996).

3.2. Pollen Identification, Abundance, and Taxonomic Diversity

Table 1. Count of main pollen types in samples from the Cameroonian continental shelf.

Family	Level	0-10	10-20	20-30	30-40	40-50	50-60	60-70	70-80	80-90	90-100	100-110	110-120	120-130	130-140	140-150	150-160	160-170	170-180	180-190	190-200	200-210	210-220	220-230	230-240	240-250	250-260	260-270	270-280	280-290	290-300	300-310	310-320	320-330	330-340	340-350	350-360	360-370	370-380	380-390	390-400	400-410	410-420	420-430	430-440	440-450	450-460	460-470	470-480	480-490	490-500	500-510	510-520	520-530	530-540	Total	Occurrence	Environments	Type																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																										
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Table 2. Pollen count data (absolute values).

Family	Taxa(Age (C14)	LEVELS (Depth)																Total																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																								
		CC15 0 - 10	20-20	30-30	40-40	50-50	60-60	70-70	80-80	90-90	100-100	110-110	120-120	130-130	140-140	150-150	160-160		170-170	180-180	190-190	200-200	210-210	220-220	230-230	240-240	250-250	260-260	270-270	280-280	290-290	300-300	310-310	320-320	330-330	340-340	350-350	360-360	370-370	380-380	390-390	400-400	410-410	420-420	430-430	440-440	450-450	460-460	470-470	480-480	490-490	500-500	510-510	520-520	530-530	540-540	550-550	560-560	570-570	580-580	590-590	600-600	610-610	620-620	630-630	640-640	650-650	660-660	670-670	680-680	690-690	700-700	710-710	720-720	730-730	740-740	750-750	760-760	770-770	780-780	790-790	800-800	810-810	820-820	830-830	840-840	850-850	860-860	870-870	880-880	890-890	900-900	910-910	920-920	930-930	940-940	950-950	960-960	970-970	980-980	990-990	1000-1000	1010-1010	1020-1020	1030-1030	1040-1040	1050-1050	1060-1060	1070-1070	1080-1080	1090-1090	1100-1100	1110-1110	1120-1120	1130-1130	1140-1140	1150-1150	1160-1160	1170-1170	1180-1180	1190-1190	1200-1200	1210-1210	1220-1220	1230-1230	1240-1240	1250-1250	1260-1260	1270-1270	1280-1280	1290-1290	1300-1300	1310-1310	1320-1320	1330-1330	1340-1340	1350-1350	1360-1360	1370-1370	1380-1380	1390-1390	1400-1400	1410-1410	1420-1420	1430-1430	1440-1440	1450-1450	1460-1460	1470-1470	1480-1480	1490-1490	1500-1500	1510-1510	1520-1520	1530-1530	1540-1540	1550-1550	1560-1560	1570-1570	1580-1580	1590-1590	1600-1600	1610-1610	1620-1620	1630-1630	1640-1640	1650-1650	1660-1660	1670-1670	1680-1680	1690-1690	1700-1700	1710-1710	1720-1720	1730-1730	1740-1740	1750-1750	1760-1760	1770-1770	1780-1780	1790-1790	1800-1800	1810-1810	1820-1820	1830-1830	1840-1840	1850-1850	1860-1860	1870-1870	1880-1880	1890-1890	1900-1900	1910-1910	1920-1920	1930-1930	1940-1940	1950-1950	1960-1960	1970-1970	1980-1980	1990-1990	2000-2000	2010-2010	2020-2020	2030-2030	2040-2040	2050-2050	2060-2060	2070-2070	2080-2080	2090-2090	2100-2100	2110-2110	2120-2120	2130-2130	2140-2140	2150-2150	2160-2160	2170-2170	2180-2180	2190-2190	2200-2200	2210-2210	2220-2220	2230-2230	2240-2240	2250-2250	2260-2260	2270-2270	2280-2280	2290-2290	2300-2300	2310-2310	2320-2320	2330-2330	2340-2340	2350-2350	2360-2360	2370-2370	2380-2380	2390-2390	2400-2400	2410-2410	2420-2420	2430-2430	2440-2440	2450-2450	2460-2460	2470-2470	2480-2480	2490-2490	2500-2500	2510-2510	2520-2520	2530-2530	2540-2540	2550-2550	2560-2560	2570-2570	2580-2580	2590-2590	2600-2600	2610-2610	2620-2620	2630-2630	2640-2640	2650-2650	2660-2660	2670-2670	2680-2680	2690-2690	2700-2700	2710-2710	2720-2720	2730-2730	2740-2740	2750-2750	2760-2760	2770-2770	2780-2780	2790-2790	2800-2800	2810-2810	2820-2820	2830-2830	2840-2840	2850-2850	2860-2860	2870-2870	2880-2880	2890-2890	2900-2900	2910-2910	2920-2920	2930-2930	2940-2940	2950-2950	2960-2960	2970-2970	2980-2980	2990-2990	3000-3000	3010-3010	3020-3020	3030-3030	3040-3040	3050-3050	3060-3060	3070-3070	3080-3080	3090-3090	3100-3100	3110-3110	3120-3120	3130-3130	3140-3140	3150-3150	3160-3160	3170-3170	3180-3180	3190-3190	3200-3200	3210-3210	3220-3220	3230-3230	3240-3240	3250-3250	3260-3260	3270-3270	3280-3280	3290-3290	3300-3300	3310-3310	3320-3320	3330-3330	3340-3340	3350-3350	3360-3360	3370-3370	3380-3380	3390-3390	3400-3400	3410-3410	3420-3420	3430-3430	3440-3440	3450-3450	3460-3460	3470-3470	3480-3480	3490-3490	3500-3500	3510-3510	3520-3520	3530-3530	3540-3540	3550-3550	3560-3560	3570-3570	3580-3580	3590-3590	3600-3600	3610-3610	3620-3620	3630-3630	3640-3640	3650-3650	3660-3660	3670-3670	3680-3680	3690-3690	3700-3700	3710-3710	3720-3720	3730-3730	3740-3740	3750-3750	3760-3760	3770-3770	3780-3780	3790-3790	3800-3800	3810-3810	3820-3820	3830-3830	3840-3840	3850-3850	3860-3860	3870-3870	3880-3880	3890-3890	3900-3900	3910-3910	3920-3920	3930-3930	3940-3940	3950-3950	3960-3960	3970-3970	3980-3980	3990-3990	4000-4000	4010-4010	4020-4020	4030-4030	4040-4040	4050-4050	4060-4060	4070-4070	4080-4080	4090-4090	4100-4100	4110-4110	4120-4120	4130-4130	4140-4140	4150-4150	4160-4160	4170-4170	4180-4180	4190-4190	4200-4200	4210-4210	4220-4220	4230-4230	4240-4240	4250-4250	4260-4260	4270-4270	4280-4280	4290-4290	4300-4300	4310-4310	4320-4320	4330-4330	4340-4340	4350-4350	4360-4360	4370-4370	4380-4380	4390-4390	4400-4400	4410-4410	4420-4420	4430-4430	4440-4440	4450-4450	4460-4460	4470-4470	4480-4480	4490-4490	4500-4500	4510-4510	4520-4520	4530-4530	4540-4540	4550-4550	4560-4560	4570-4570	4580-4580	4590-4590	4600-4600	4610-4610	4620-4620	4630-4630	4640-4640	4650-4650	4660-4660	4670-4670	4680-4680	4690-4690	4700-4700	4710-4710	4720-4720	4730-4730	4740-4740	4750-4750	4760-4760	4770-4770	4780-4780	4790-4790	4800-4800	4810-4810	4820-4820	4830-4830	4840-4840	4850-4850	4860-4860	4870-4870	4880-4880	4890-4890	4900-4900	4910-4910	4920-4920	4930-4930	4940-4940	4950-4950	4960-4960	4970-4970	4980-4980	4990-4990	5000-5000	5010-5010	5020-5020	5030-5030	5040-5040	5050-5050	5060-5060	5070-5070	5080-5080	5090-5090	5100-5100	5110-5110	5120-5120	5130-5130	5140-5140	5150-5150	5160-5160	5170-5170	5180-5180	5190-5190	5200-5200	5210-5210	5220-5220	5230-5230	5240-5240	5250-5250	5260-5260	5270-5270	5280-5280	5290-5290	5300-5300	5310-5310	5320-5320	5330-5330	5340-5340	5350-5350	5360-5360	5370-5370	5380-5380	5390-5390	5400-5400	5410-5410	5420-5420	5430-5430	5440-5440	5450-5450	5460-5460	5470-5470	5480-5480	5490-5490	5500-5500	5510-5510	5520-5520	5530-5530	5540-5540	5550-5550	5560-5560	5570-5570	5580-5580	5590-5590	5600-5600	5610-5610	5620-5620	5630-5630	5640-5640	5650-5650	5660-5660	5670-5670	5680-5680	5690-5690	5700-5700	5710-5710	5720-5720	5730-5730	5740-5740	5750-5750	5760-5760	5770-5770	5780-5780	5790-5790	5800-5800	5810-5810	5820-5820	5830-5830	5840-5840	5850-5850	5860-5860	5870-5870	5880-5880	5890-5890	5900-5900	5910-5910	5920-5920	5930-5930	5940-5940	5950-5950	5960-5960	5970-5970	5980-5980	5990-5990	6000-6000	6010-6010	6020-6020	6030-6030	6040-6040	6050-6050	6060-6060	6070-6070	6080-6080	6090-6090	6100-6100	6110-6110	6120-6120	6130-6130	6140-6140	6150-6150	6160-6160	6170-6170	6180-6180	6190-6190	6200-6200	6210-6210	6220-6220	6230-6230	6240-6240	6250-6250	6260-6260	6270-6270	6280-6280	6290-6290	6300-6300	6310-6310	6320-6320	6330-6330	6340-6340	6350-6350	6360-6360	6370-6370	6380-6380	6390-6390	6400-6400	6410-6410	6420-6420	6430-6430	6440-6440	6450-6450	6460-6460	6470-6470	6480-6480	6490-6490	6500-6500	6510-6510	6520-6520	6530-6530	6540-6540	6550-6550	6560-6560	6570-6570	6580-6580	6590-6590	6600-6600	6610-6610	6620-6620	6630-6630	6640-6640	6650-6650	6660-6660	6670-6670	6680-6680	6690-6690	6700-6700	6710-6710	6720-6720	6730-6730	6740-6740	6750-6750	6760-6760	6770-6770	6780-6780	6790-6790	6800-6800	6810-6810	6820-6820	6830-6830	6840-6840	6850-6850	6860-6860	6870-6870	6880-6880	6890-6890	6900-6900	6910-6910	6920-6920	6930-6930	6940-6940	6950-6950	6960-6960	6970-6970	6980-6980	6990-6990	7000-7000	7010-7010	7020-7020	7030-7030	7040-7040	7050-7050	7060-7060	7070-7070	7080-7080	7090-7090	7100-7100	7110-7110	7120-7120	7130-7130	7140-7140	7150-7150	7160-7160	7170-7170	7180-7180	7190-7190	7200-7200	7210-7210	7220-7220	7230-7230	7240-7240	7250-7250	7260-7260	7270-7270	7280-7280	7290-7290	7300-7300	7310-7310	7320-7320	7330-7330	7340-7340	7350-7350	7360-7360	7370-7370	7380-7380	7390-7390	7400-7400	7410-7410	7420-7420	7430-7430	7440-7440	7450-7450	7460-7460	7470-7470	7480-7480	7490-7490	7500-7500	7510-7510	7520-7520	7530-7530	7540-7540	7550-7550	7560-7560	7570-7570	7580-7580	7590-7590	7600-7600	7610-7610	7620-7620	7630-7630	7640-7640	7650-7650	7660-7660	7670-7670	7680-7680	7690-7690	7700-7700	7710-7710	7720-7720	7730-7730	7740-7740	7750-7750	7760-7760	7770-7770	7780-7780	7790-7790	7800-7800	7810-7810	7820-7820	7830-7830	7840-7840	7850-7850	7860-7860	7870-7870	7880-7880	7890-7890	7900-7900	7910-7910	7920-7920	7930-7930	7940-7940	7950-7950	7960-7960	7970-7970	7980-7980	7990-7990	8000-8000	8010-8010	8020-8020	8030-8030	8040-8040	8050-8050	8060-8060	8070-8070	8080-8080	8090-8090	8100-8100	8110-8110	8120-8120	8130-8130	8140-8140	8150-8150	8160-8160	8170-8170	8180-8180	8190-8190	8200-8200	8210-8210	8220-8220	8230-8230	8240-8240

Examining three to four transects per slide was generally sufficient to reach the required 300 pollen grains, except at the 70 - 80 cm level (seven transects) and the 430 - 440 cm level (two transects). In total, 26,444 pollen and spore grains were identified. Among other palynomorphs, approximately 939 were indeterminate, 132 were damaged, and 785 corresponded to dinoflagellate cysts (**Table 1**). The 224 recorded taxa, distributed among 93 families, show a marked dominance of *Rhizophora* (34.3%), spores (18.3%), *Poaceae* (10.1%), *Cyperaceae* (8.2%), and *Alchornea* (4.7%). The highest taxonomic diversity was observed in Rubiaceae (23 genera), followed by Euphorbiaceae (16) and Phyllanthaceae (11). These assemblages reflect the regional mosaic of forested and open vegetation types.

3.3. Pollen Diagram of the Main Taxa

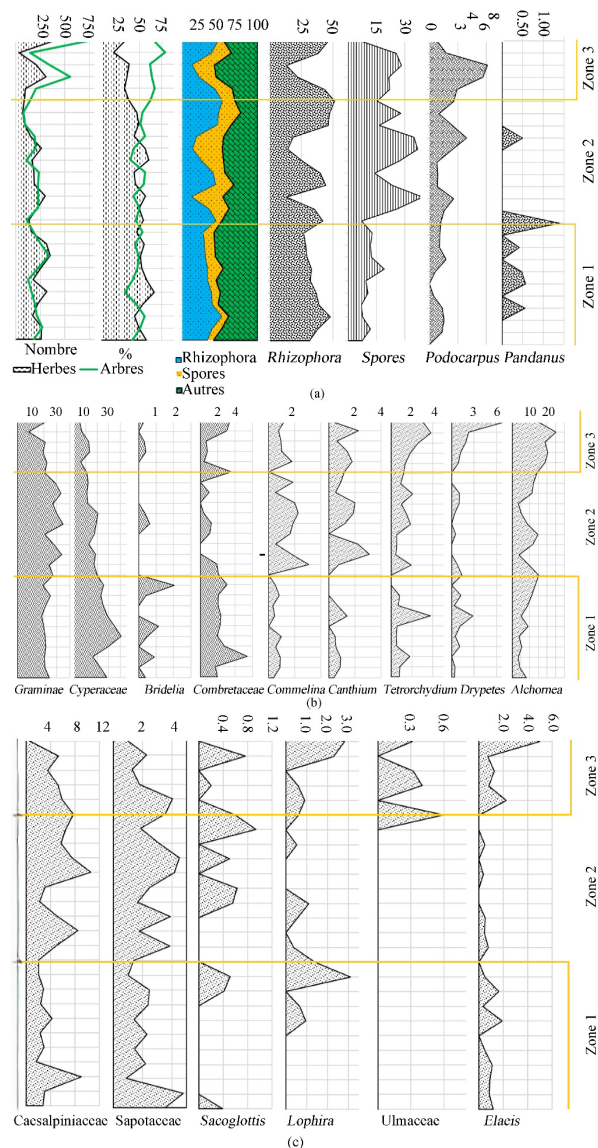
The pollen spectra forming the overall pollen diagram were constructed using two calculation matrices. The bioindicator taxa, which represent the two main pollen groups, were derived from matrix (1), which includes all identified pollen types. Matrix (2), used for the analysis of the other major and characteristic taxa, was generated from a dataset in which *Rhizophora* and spore taxa were excluded (**Table 2**).

Based on the temporal comparison of pollen spectra, significant changes in floristic composition were identified, allowing the distinction of three major successive palynological zones (**Table 3**), which correspond to the main paleoecological phases recorded in core C61.

Table 3. Description of the palynological zones.

		Zone 1	Zone 2	Zone 3
Ages	<i>Start</i>	11,400	9400	5400
	<i>End</i>	9400	5400	1450
Depth	<i>Start</i>	540	180	60
	<i>End</i>	190	60	0
Périod (years)	-	2.000	4.000	4.000
Thickness (cm)	-	350	120	60
Sedimentation rate	-	175	30	15
Séries	-	Extensive	Less condensed	Condensed
Plants	-	Shrub	Herbs	Trees
Environments	-	Mangrove	Grassy savanna	Afromontane Forest
Flora	-	<i>Rhizophora</i>	Graminae	<i>Podocarpus</i>
Spores	-	Low	Significant	Medium
<i>Podocarpus</i>	-	Low	Medium	Significant
<i>Rhizophora</i>	-	Significant	Contrast with Spores and Podocarpus	

The pollen groupings (Table 4), organized by plant type within the three palynological zones of the overall pollen diagram (Figure 4), are characterized by the following main plant assemblages: 1) biological forms (herbaceous and arboreal types); 2) taxa dominant in pollen proportions (*Rhizophora*, spores); 3) taxa serving as environmental and climatic bioindicators (*Rhizophora*, *Pandanus*, *Podocarpus*); and 4) taxa characteristic of the major ecosystems—savanna (*Poaceae*, *Bridelia*, *Combretum*), dense humid forest (*Caesalpinaceae*, *Sapotaceae*, *Ulmaceae*, *Sacoglottis*), and ubiquitous or pioneer understory taxa (*Alchornea*, *Tetrachidium*, *Drypetes*).



Zone 1: Regression of ubiquitous pioneer and undergrowth taxa (9400 - 11300 years). **Zone 2:** Increase in spores during the Podocarpus phase, both alternating in contrast with *Rhizophora* (5400 - 9400 years). **Zone 3:** Forest resurgence in favor of Podocarpus (0 - 5400 years).

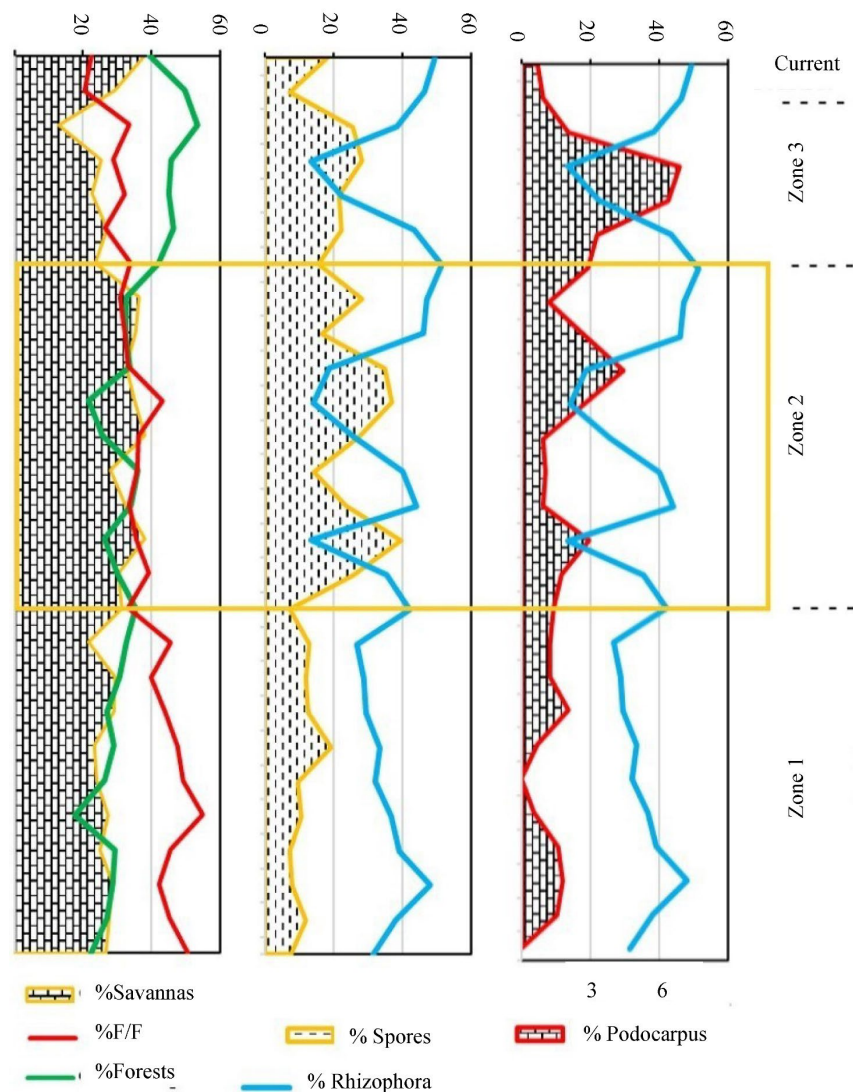
Figure 4. Pollen diagram of ecological bioindicators.

3.4. Spectra of Paleoecological Bioindicators

Table 5. Vegetation types and ecological bioindicators.

	CC135	0 - 10	10 - 20	20 - 30	30 - 40	40 - 50	50 - 60	60 - 70	70 - 80	80 - 90	90 - 100	100 - 110	110 - 120	120 - 130	130 - 140	140 - 150	150 - 180	180 - 190	190 - 220	220 - 230	230 - 260	260 - 270	270 - 300	300 - 310	310 - 330	330 - 340	340 - 350	350 - 360	360 - 390	390 - 400	400 - 430	430 - 440	440 - 510	510 - 520	520 - 540	540 - 590	Total						
Forests	142	532	98	249	389	144	99	49	62	117	100	72	152	142	138	137	89	89	112	173	187	187	142	81	84	84	124	124	112	145	145	100	3968										
F/S	82	221	62	158	277	83	80	46	62	116	200	102	152	140	187	179	84	156	226	303	303	233	151	261	191	261	191	165	243	243	226	4381											
Savannas	138	315	24	138	196	84	56	55	68	114	163	107	119	135	198	138	79	74	166	201	114	74	131	105	112	149	119	3370															
Total	362	1068	183	544	861	311	235	150	192	347	462	280	422	418	523	453	252	341	565	691	490	305	476	420	389	537	445	445	445	445	445	445	445	445	445	445	445	445	445	445			
%Forests	39	50	53	46	45	46	42	33	32	34	22	26	36	34	26	30	35	33	33	31	27	29	26	18	29	29	27	22	22	33													
%F/S	23	21	34	29	32	27	34	31	32	33	43	36	36	34	36	39	33	46	40	44	48	49	55	46	42	45	51	38															
%Savannas	38	30	13	25	23	27	24	36	35	33	35	38	28	32	38	31	31	22	29	29	23	24	27	25	29	28	27	29															
%Total	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100	100
<i>Rhizophora</i> (%)	49.61	###	38.36	13.33	22.40	43.40	51.55	47.33	46.29	18.82	14.24	25.89	40.14	44.33	13.37	35.55	42.25	26.80	28.70	29.57	33.33	32.26	36.96	39.11	48.11	38.30	31.48	###															
Spores %	18.11	7.16	25.58	28.25	21.75	22.14	15.64	28.19	16.61	35.01	36.95	27.25	14.22	23.52	39.52	25.64	7.29	12.98	11.96	12.54	19.38	9.68	10.66	7.26	7.86	11.99	7.71	###															
<i>Podocarpus</i> %	0.59	0.81	1.79	6.14	5.68	2.93	2.58	1.03	2.47	3.94	2.37	0.82	0.92	0.85	2.59	1.56	1.22	1.10	1.09	1.79	0.63	0.00	0.45	1.40	1.57	1.35	0.00	1.77															

The bioindicator diagram, constructed from the data presented in **Table 5** and illustrated in **Figure 5**, reveals marked variations through time. Alternating fluctuations in intensity are observed between, on the one hand, plant formations (forested, savanna, and mixed types) and, on the other hand, specific taxa such as *Rhizophora*, spores, and *Podocarpus*. *Rhizophora* pollen is particularly abundant in Zone 1, whereas spores and *Podocarpus* are only sparsely represented. In Zone 2, the proportion of spores increases significantly, indicating a more open and less humid phase. Finally, in Zone 3, *Podocarpus* exhibits a steady increase toward its maximum abundance, reflecting the expansion of montane or temperate vegetation under cooler and more humid climatic conditions.



Zone 1: Regression of ubiquitous pioneer and undergrowth taxa (9400 - 11,300 years). **Zone 2:** Increase in spores during the Podocarpus phase, both alternating in contrast with *Rhizophora* (5400 - 9400 years). **Zone 3:** Forest resurgence in favor of *Podocarpus* (0 - 5400 years)

Figure 5. Evolution of ecological bioindicators (*Rhizophora*, *Podocarpus*, Spores).

4. Discussion

4.1. Taxonomic Diversity of Plant Families

Pollen analysis reveals a substantial representation of numerous plant families, indicating high diversity at both the genus and species levels. The most abundant taxa, in terms of both absolute number and relative proportion, correspond to those identified in recent dredging samples from the Cameroonian continental shelf, notably *Rhizophora*, spores, *Poaceae*, *Cyperaceae*, *Alchornea*, *Drypetes*, *Uapaca*, and *Podocarpus* (Bengo et al., 2020; Bengo et al., 2025). These results corroborate previous observations on marine pollen sedimentation, emphasizing the key role of fluvial inputs, particularly from the Sanaga River—the main drainage system collecting runoff from watersheds that extend from the dense forests of the south to the savannas of central and northern Cameroon (Olivry, 1986; Girresse et al., 1995). Thus, the floristic diversity recorded in core C61 faithfully reflects the botanical richness of the adjacent continent.

However, overall taxonomic diversity remains slightly lower than that observed on land, likely due to taphonomic processes and the constraints imposed by marine dynamics (Fredoux, 1994; Dupont & Agwu, 1991; Vincens et al., 1994). Indeed, incomplete fossilization and the long-distance transport of pollen grains from their continental source can result in the underrepresentation or even absence of certain taxa originating from distant ecosystems. In particular, arboreal taxa characteristic of savanna or semi-deciduous forest, montane or semi-deciduous forests—such *Hymenocardia*, *Bridelia* or *Annona*—are often poorly represented or absent in marine records, despite being common in the inland vegetation. This limitation, well documented in marine palynological studies from Central Africa and the Gulf of Guinea (Caratini et al., 1987; Van Campo & Bengo, 2004; Hernandez et al., 2021), may affect the accuracy of paleoenvironmental reconstructions based solely on marine records.

4.2. Holocene Paleoenvironmental Evolution According to Core C61 (Figure 4)

The palynological data from core C61 allow for a detailed reconstruction of paleoenvironmental dynamics on the Cameroonian continental shelf from the onset of the Holocene to approximately 1450 cal yr BP. Rather than beginning with general considerations, we first present the main vegetational and sedimentary stages identified in the core and interpret them in light of regional climatic trends and sea-level changes. Three distinct paleoecological phases are identified, reflecting the evolving balance between coastal dynamics, continental inputs, and climatic forcing.

a) First Phase (11,400 - 9400 cal yr BP): Coastal Expansion and Fluvio-lagoonal Environments Dominated by *Rhizophora* and *Cyperaceae*

This initial phase, corresponding to Zone 1, is characterized by the dominance of *Rhizophora* and *Cyperaceae*, with a co-occurrence of taxa such as *Combretaceae*, *Bridelia*, and various herbaceous pollen types (Figure 4). These

assemblages indicate the development of mangrove systems and wetland savannas at the land–sea interface, under the influence of a warm and humid climate. The high abundance of *Rhizophora* suggests a well-established mangrove belt, while *Cyperaceae* and NAP taxa reflect seasonally waterlogged open habitats. *Podocarpus* is poorly represented, likely restricted to mid-altitude zones due to still-elevated regional temperatures (Maley, 1987; Lebamba et al., 2012).

This period follows the Late Dryas and corresponds to a time of increasing monsoon strength and hydrological intensification in Central Africa. Enhanced runoff from the Sanaga watershed, possibly linked to glacial meltwater inputs, likely contributed to accelerated sedimentation on the continental shelf (Giresse et al., 1995; Ngueutchoua, 1996). However, the role of glacial melting as a climatic driver must be interpreted with caution in this tropical basin, where local hydrodynamics and ITCZ displacement are more significant.

b) Second Phase (9400 - 5400 cal yr BP): Stabilization of Coastal Systems and Expansion of Savanna-Grassland Mosaics

During this interval, *Cyperaceae* decline while *Poaceae* and spores increase, accompanied by a retreat of shrubs such as *Bridelia* and *Combretaceae*. This floral transition suggests a shift toward more open environments, such as wooded savannas or anthropogenic clearings, but does not unequivocally signal aridification. Contrary to earlier interpretations, the available data do not support a markedly dry climate during this phase.

Instead, this period likely corresponds to the Holocene Climatic Optimum, characterized regionally by high rainfall and increased vegetation productivity (Ngomanda et al., 2009a, 2009b; Gasse, 2000). The expansion of grasses may reflect disturbance regimes, seasonal contrasts, or hydro-edaphic variability, rather than long-term drought. Simultaneously, a marine transgression reached its maximum around 6000 cal yr BP, gradually distancing the core site from the shoreline. This change in sedimentary context explains the dominance of marine muds and the reduced terrigenous input in this unit (Giresse et al., 1995; Ngueutchoua, 1996).

c) Third Phase (5400 - 1450 cal yr BP): Forest Expansion and Climatic Oscillations in a Stabilized Marine Context

This phase, corresponding to the upper part of the core, is marked by a sharp decrease in sedimentation rate (Figure 5), indicating post-transgressive stabilization of the marine depositional environment. The pollen assemblage reveals a diversification of arboreal taxa, including *Caesalpiniaceae*, *Sapotaceae*, and *Ulmaceae*, along with forest species such as *Sacoglottis*, *Lophira*, and *Elaeis*. Pioneer taxa (*Alchornea*, *Tetrorchidium*, *Pycnanthus*, *Uapaca*, *Drypetes*) are also well represented, pointing to mature forest development with dynamic understorey regeneration.

The occurrence of *Podocarpus*, coupled with peak values of spores, indicates episodes of cooler and wetter conditions, possibly linked to cloud-based occult

precipitation in upland areas (Maley, 1987). Rather than a simple trend, this period reflects alternating wet and dry subphases, consistent with regional climatic variability during the Late Holocene (Lebamba et al., 2012; Suchel, 1988; Ngomanda et al., 2009a, 2009b). The observed vegetational transitions support a scenario of increased forest cover in the last two millennia, aligning with other Central African paleoecological records (Reynaud-Farrera et al., 1996; Maley et al., 2017; Yao et al., 2025).

4.3. The Importance of Bioindicators in Paleoenvironmental Interpretation

4.3.1. Variations in the Proportion of *Rhizophora* Spores and Pollen

Analyses performed on core C61 reveal that the high proportions of spores observed at various levels mainly result from a decrease in *Rhizophora* pollen. Spores are produced in large quantities by terrestrial or epiphytic ferns in a wide range of environments, from savannas to forests, including transitional zones. Such spore abundance is observed near the mouths of all coastal rivers. This interpretation is supported by modern observations from the Cameroonian continental shelf, which show that present-day concentrations display similar trends—that is, peaks in *Rhizophora* and spore abundance often coincide. However, the relative magnitudes of these peaks vary among dredged samples, and the proportions of each taxon depend on their proximity to mangrove development areas (Bengo et al., 2025).

However, a clear spatial differentiation is observed: in the southern sector, where the coastline is devoid of mangroves, spore values reach high levels, whereas in the north, the presence of tidal influence, extensive mangroves, and the high production of *Rhizophora* pollen maintain these values at lower levels.

4.3.2. Dynamics of *Rhizophora*, Spores, and the Paleoecological Significance of the Coastal Signal

The variations in *Rhizophora* pollen frequency observed in core C61 reflect changes in the extent or proximity of mangrove ecosystems during the Holocene. *Rhizophora* is an excellent indicator of estuarine and lagoonal zones under saline influence, and its abundant presence generally suggests that the recording site was located relatively close to the shoreline under a stabilized sea level (Fredoux, 1994; Bengo et al., 2025; Giresse et al., 1995). Sharp declines in *Rhizophora* within certain sections of the pollen spectrum may reflect either a gradual offshore shift of the coastline due to Holocene transgression (Ngueutchoua, 1996) or local ecological factors (e.g., sedimentation, salinity, turbidity) unfavorable to mangrove development (Dupont & Agwu, 1991).

In the same intervals, relative increases in spores probably reflect dilution effects in the pollen signal rather than an actual rise in spore production. The inverse alternation between *Rhizophora* and spores, particularly in Zones 2 and 3, therefore indicates variations in fluvial input and coastal vegetation cover, without necessarily implying a regressive or markedly cooler climatic episode (Bengo et

al., 2025; Lebamba et al., 2009).

Moreover, the intermittent presence of *Podocarpus* cannot be interpreted as direct evidence of altitudinal migration due to cooling. It is more likely that these occurrences reflect sporadic inputs from mid-altitude areas during wetter phases favorable to its regional development (Maley, 1987; Lebamba et al., 2012). In the absence of independent data on sea level or temperature, caution is warranted in interpreting these signals. Thus, the opposite variations among *Rhizophora*, spores, and *Podocarpus* should be viewed as the combined result of sedimentary dynamics, distance from the coastline, local ecosystem fluctuations, and taphonomic effects inherent to marine environments (Fredoux, 1994; Dupont & Agwu, 1991; Van Campo & Bengo, 2004).

4.4. Regional Comparison and Paleoclimatic Consistency

The palynological sequence from core C61 shows several transitions in vegetation composition that are partly comparable to continental records from Central Africa. However, some key intervals well documented in lake records—such as the forest retreat and climatic instability around 2500 cal yr BP—are not clearly expressed in this marine core.

This retreat has already been highlighted in recent syntheses of Holocene environmental change in Central Africa (Giresse & Nguetichoua, 1998; Nguetichoua & Giresse, 2010), which emphasize the widespread nature of the 2.5 ka dry phase, often associated with savanna expansion, lake-level fluctuations, and cultural changes. Its absence or attenuation in core C61 could reflect several factors: the limited number of palynological levels in the upper part of the core, the large intervals between radiocarbon dates, or taphonomic constraints linked to low sedimentation rates (Figure 3).

Despite this limitation, the overall trends observed in core C61—namely, mid-Holocene forest expansion and Late Holocene fluctuations—remain broadly consistent with other regional archives, such as Lake Barombi Mbo (Brenac, 1988), Lake Ossa (Reynaud-Farrera, 1995), and Kitina (Elenga et al., 1996). These comparisons highlight the potential of marine pollen records to capture large-scale climatic phases, but also their limits in recording short-term or abrupt events, especially when sampling resolution is low.

Future work should therefore focus on improving the age-depth model of core C61 through additional radiocarbon dates, particularly in the uppermost sections, and increasing palynological resolution between 3000 and 1000 cal yr BP to verify whether regional signals such as the 2.5 ka dry event are preserved in the marine domain.

5. Conclusion

The analysis of core C61 provides useful insights into the evolution of coastal environments along the Cameroonian margin over the past ~11,000 years. Three

major vegetation phases were identified: an initial phase dominated by *Rhizophora* and *Cyperaceae* under humid, fluviially active conditions; a second phase reflecting more open landscapes (*Poaceae*, spores); and a third phase indicating the expansion of humid forests (*Caesalpiniaceae*, *Sapotaceae*, *Podocarpus*), likely under more stable climatic conditions.

Despite these constraints, the C61 sequence confirms the potential of marine deposits to reflect long-term regional trends. It also emphasizes the need for high-resolution, multi-proxy approaches to distinguish climatic signals from sedimentary and taphonomic noise.

Conflicts of Interest

The authors declare no conflicts of interest regarding the publication of this paper.

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